

Fonte du **pergélisol** et décomposition microbienne: quels impacts sur le **cycle du carbone** ?

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Permafrost is a **huge carbon reservoir**

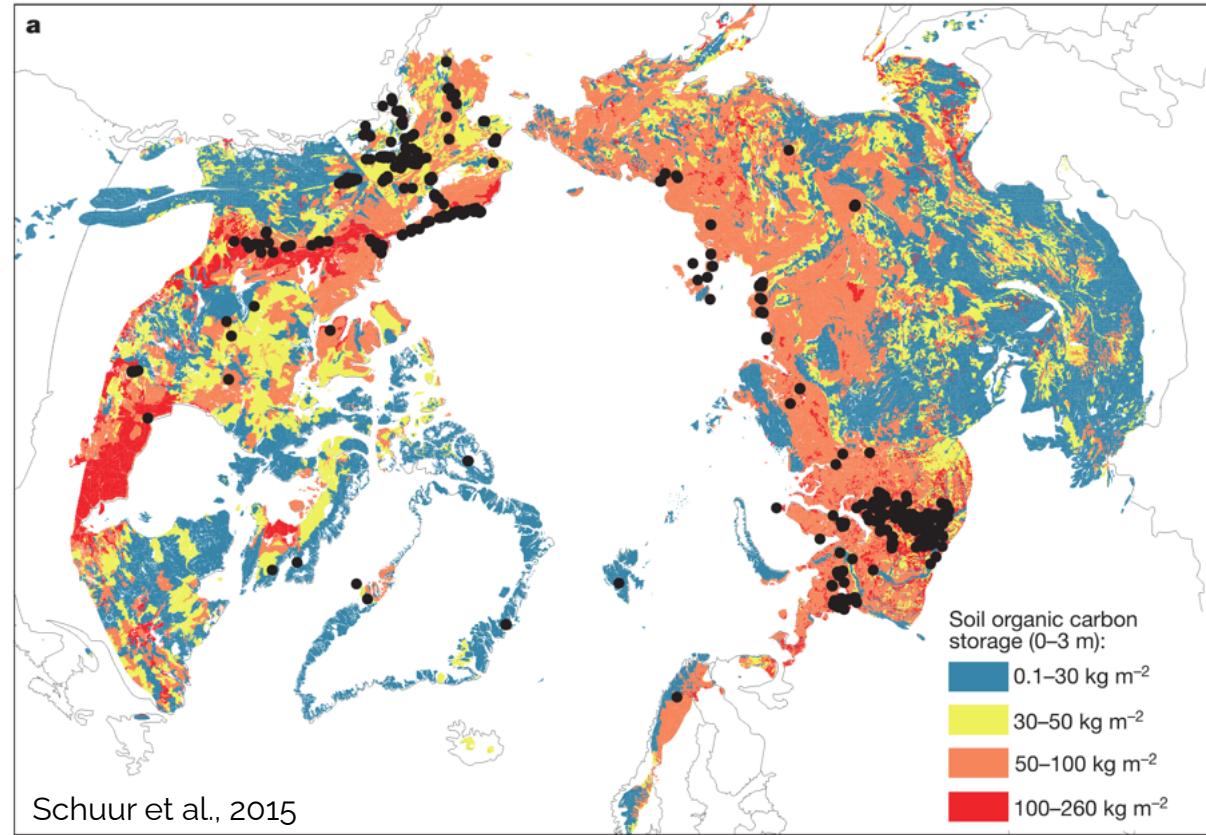
Permafrost covers **~15 million km²**

→ 11% of continental surfaces

Permafrost soils contain **~1500 PgC**

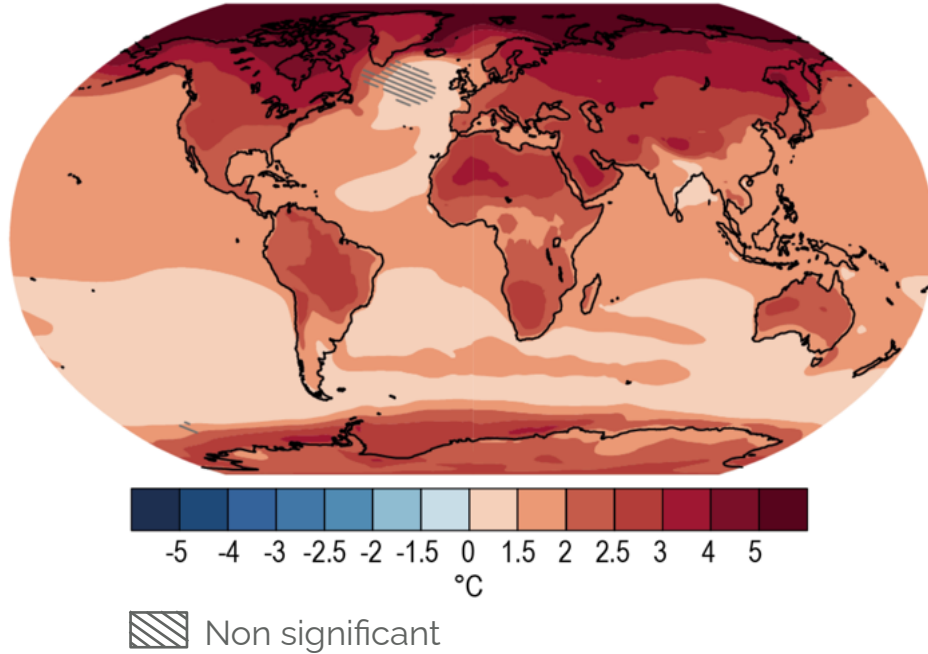
→ 2 times the atmospheric carbon content

→ 60% of the world's soil carbon



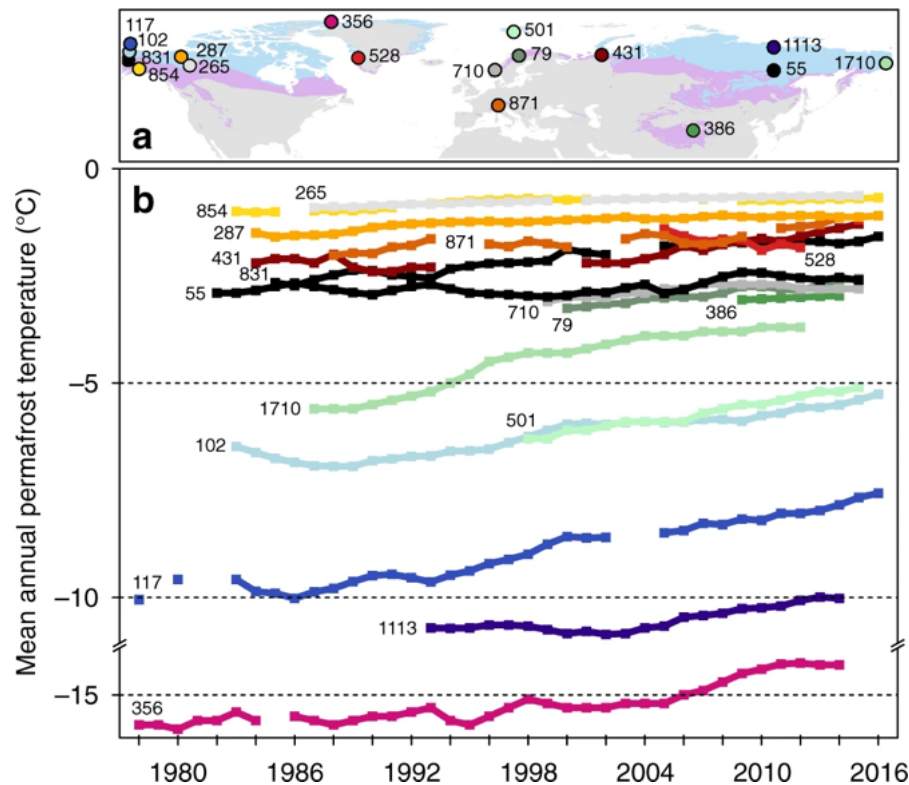
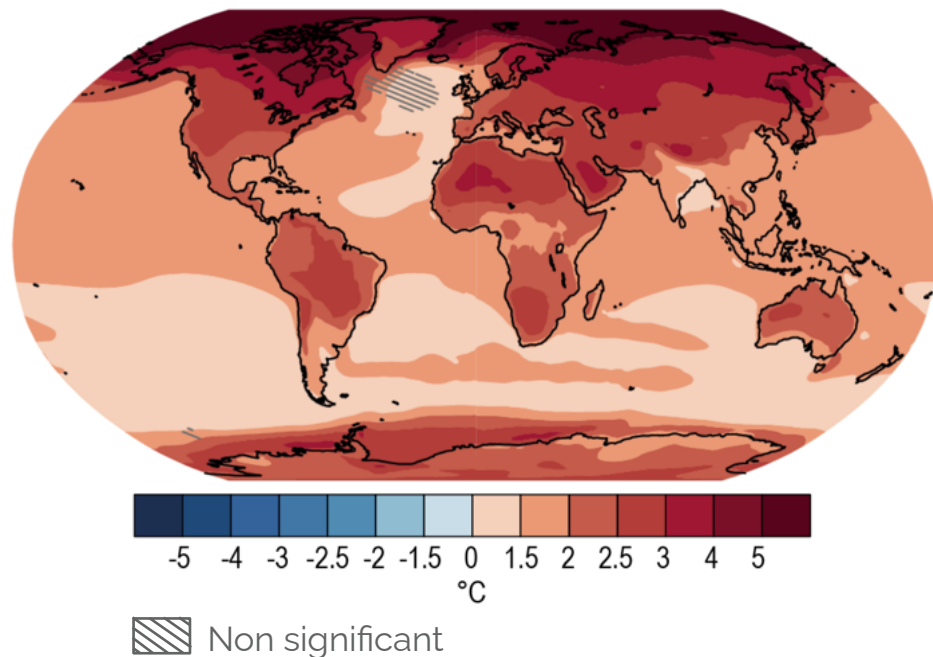
The Arctic is **warming faster** than the rest of the globe

Surface air temperature change at +2°C global warming
(model projection)

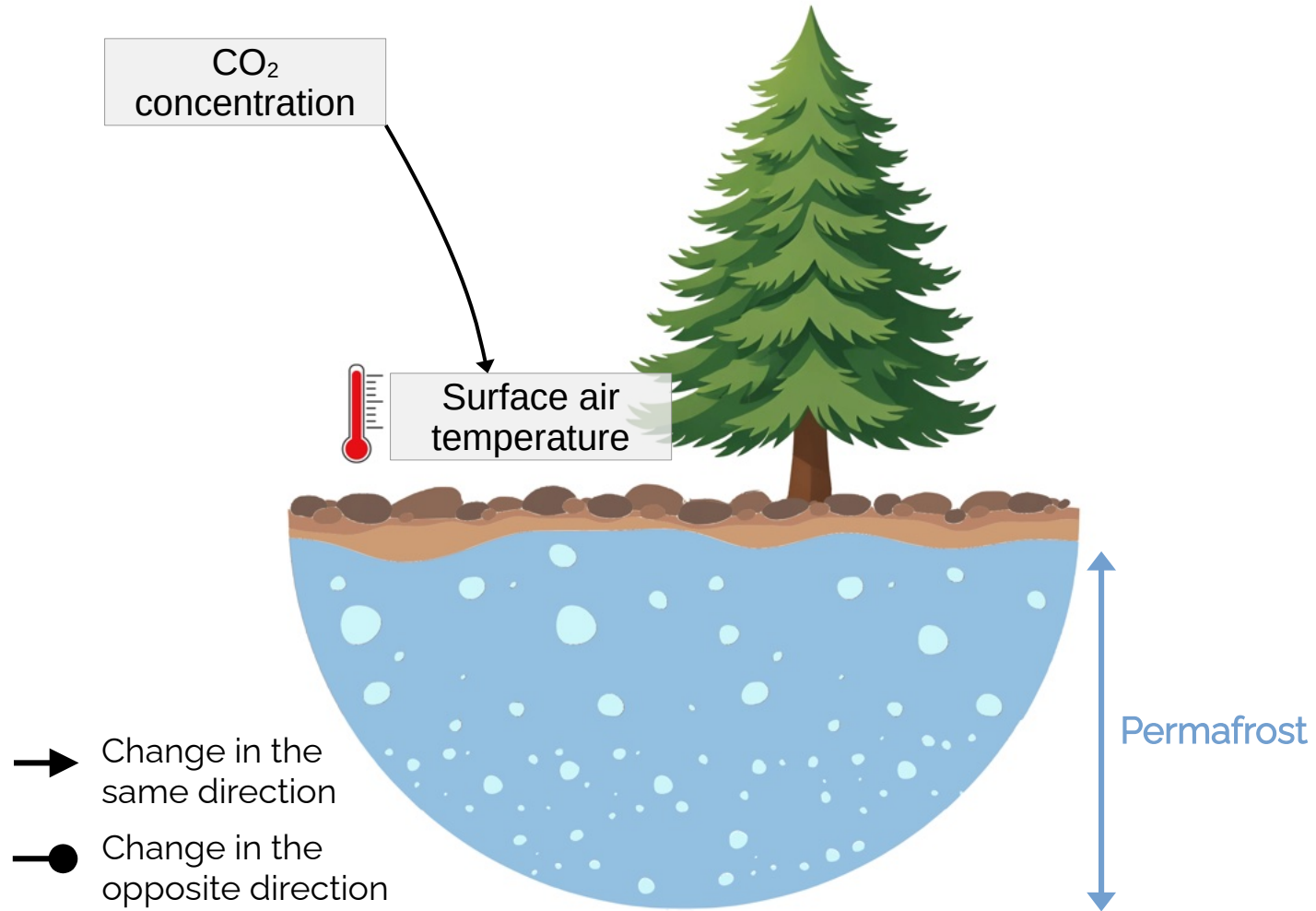


The Arctic is **warming faster** than the rest of the globe, resulting in **permafrost warming**

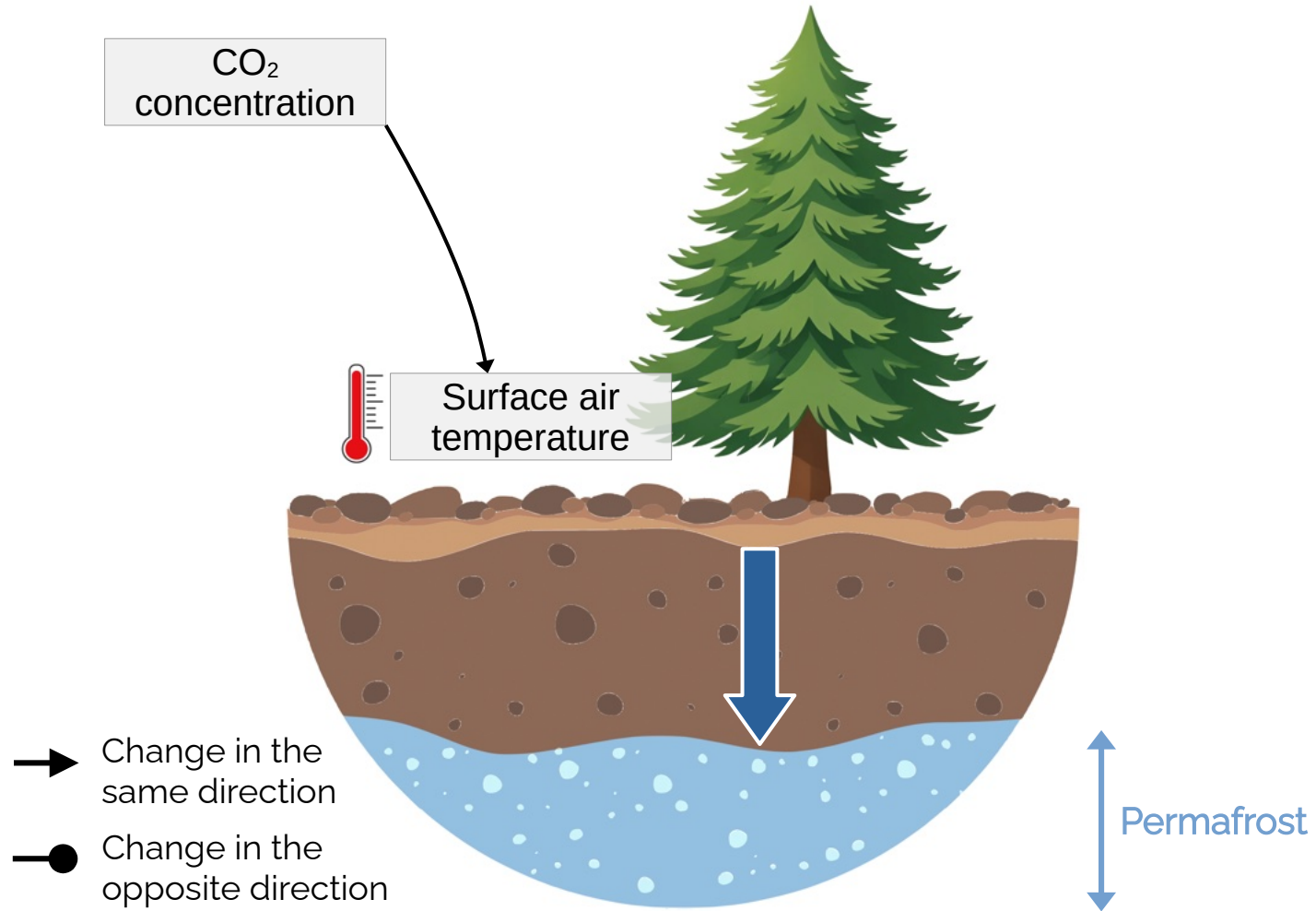
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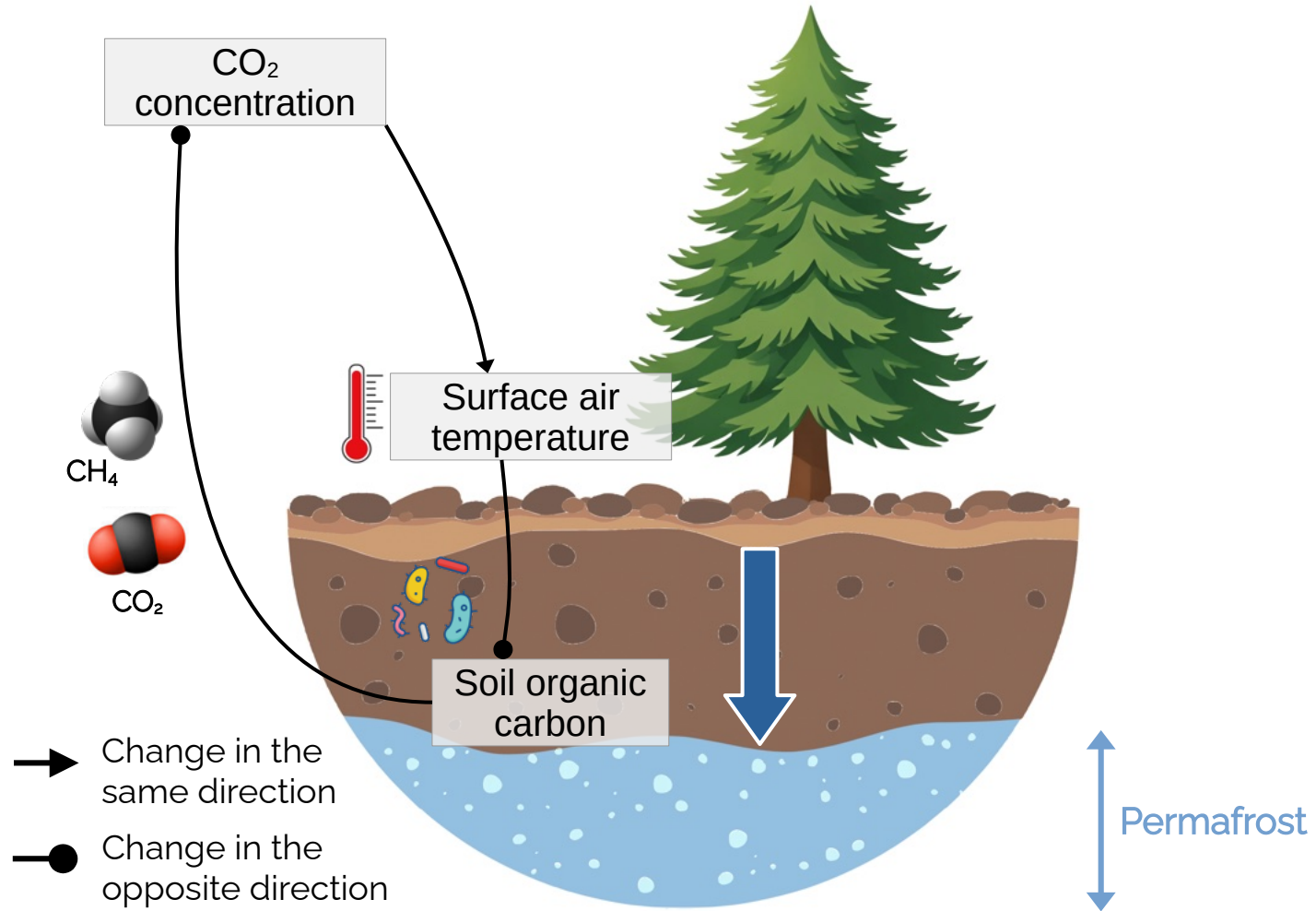
The **permafrost carbon-climate feedback** could amplify global warming



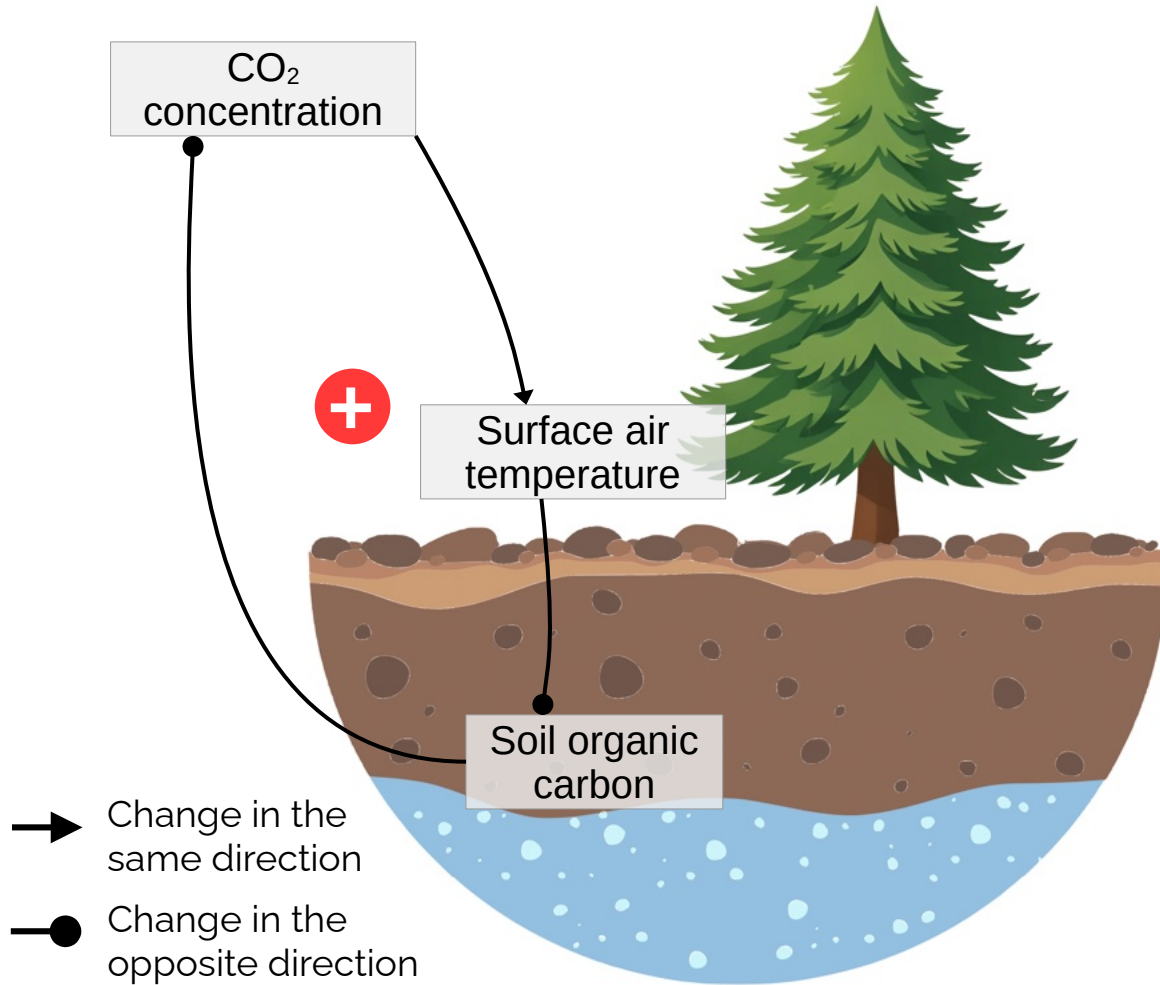
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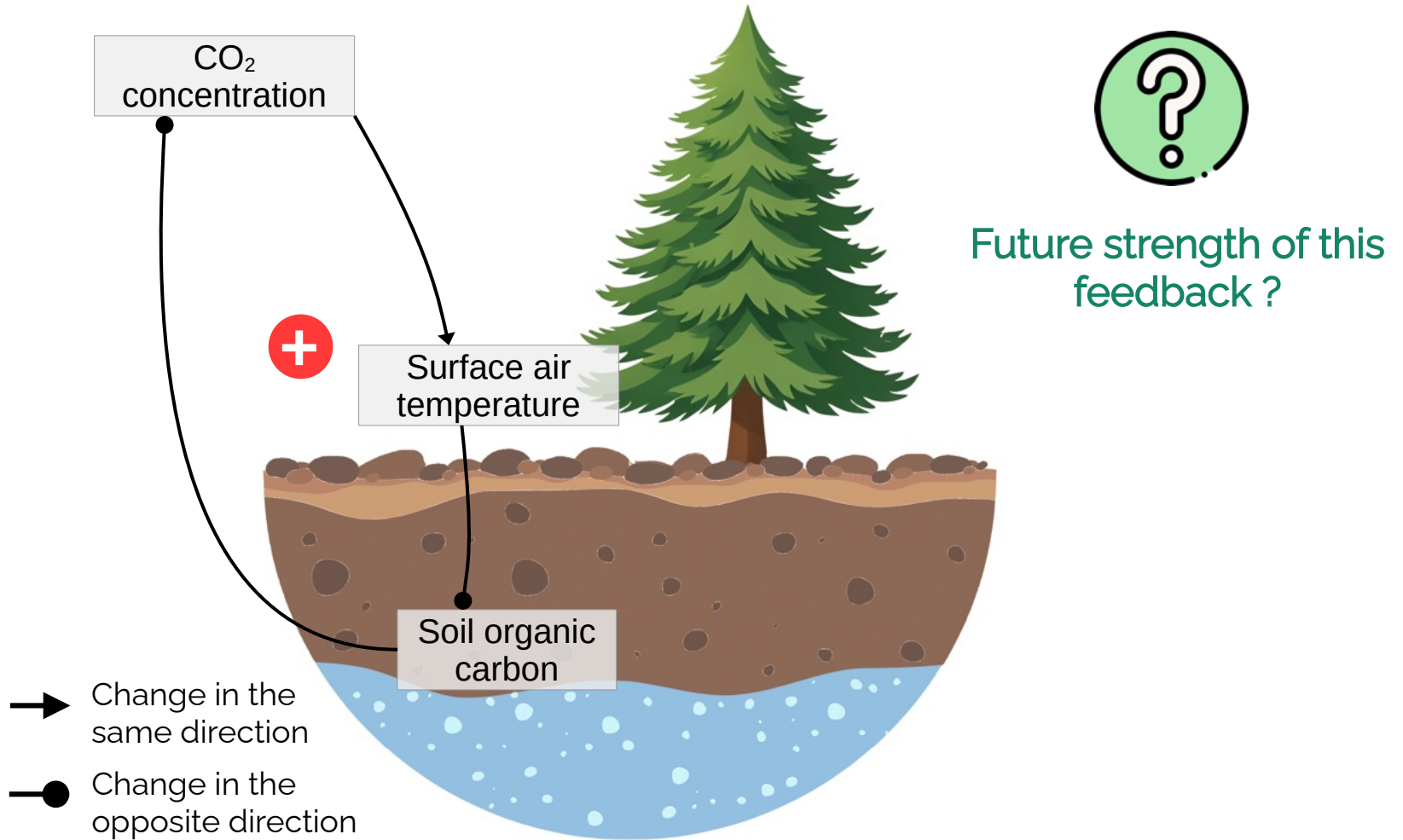
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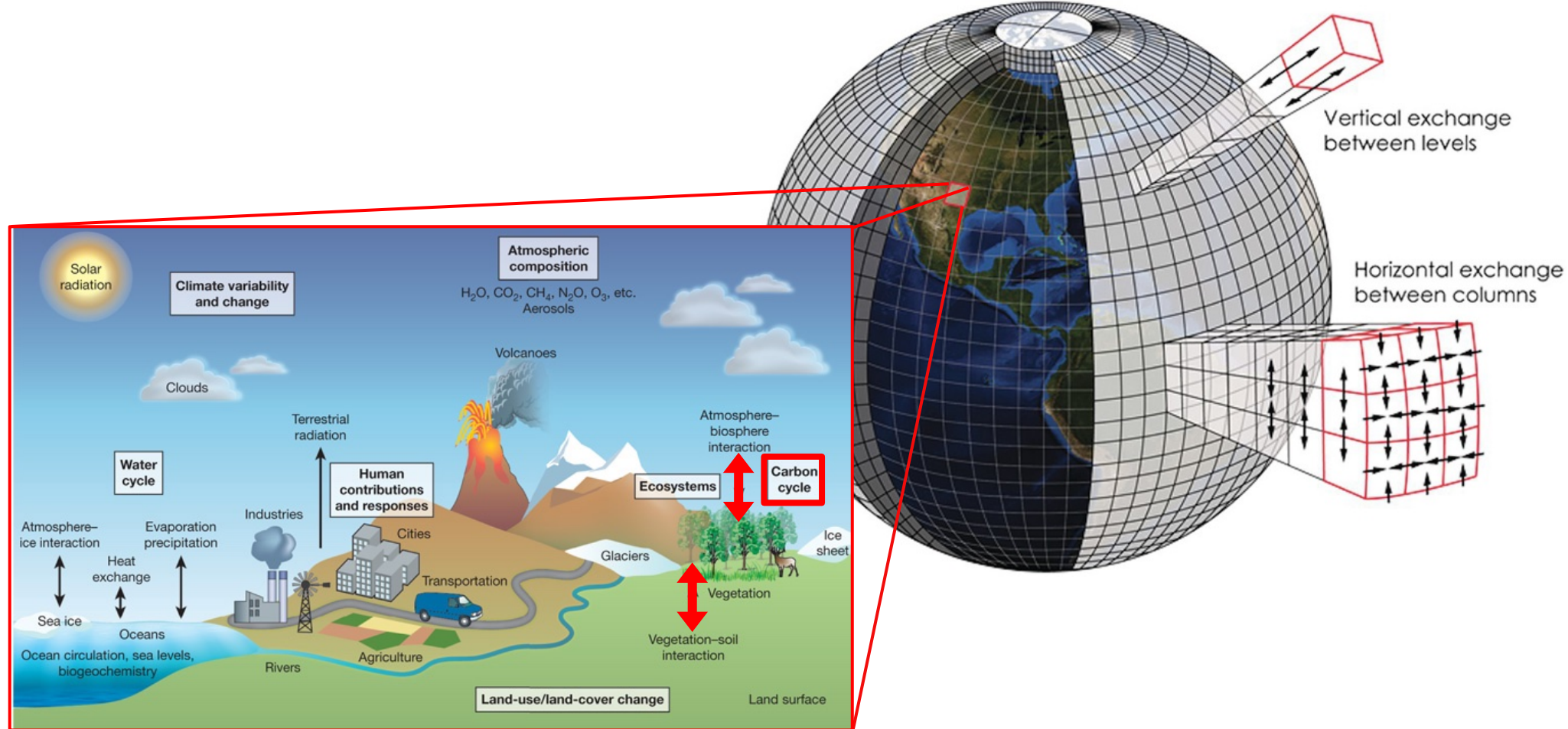
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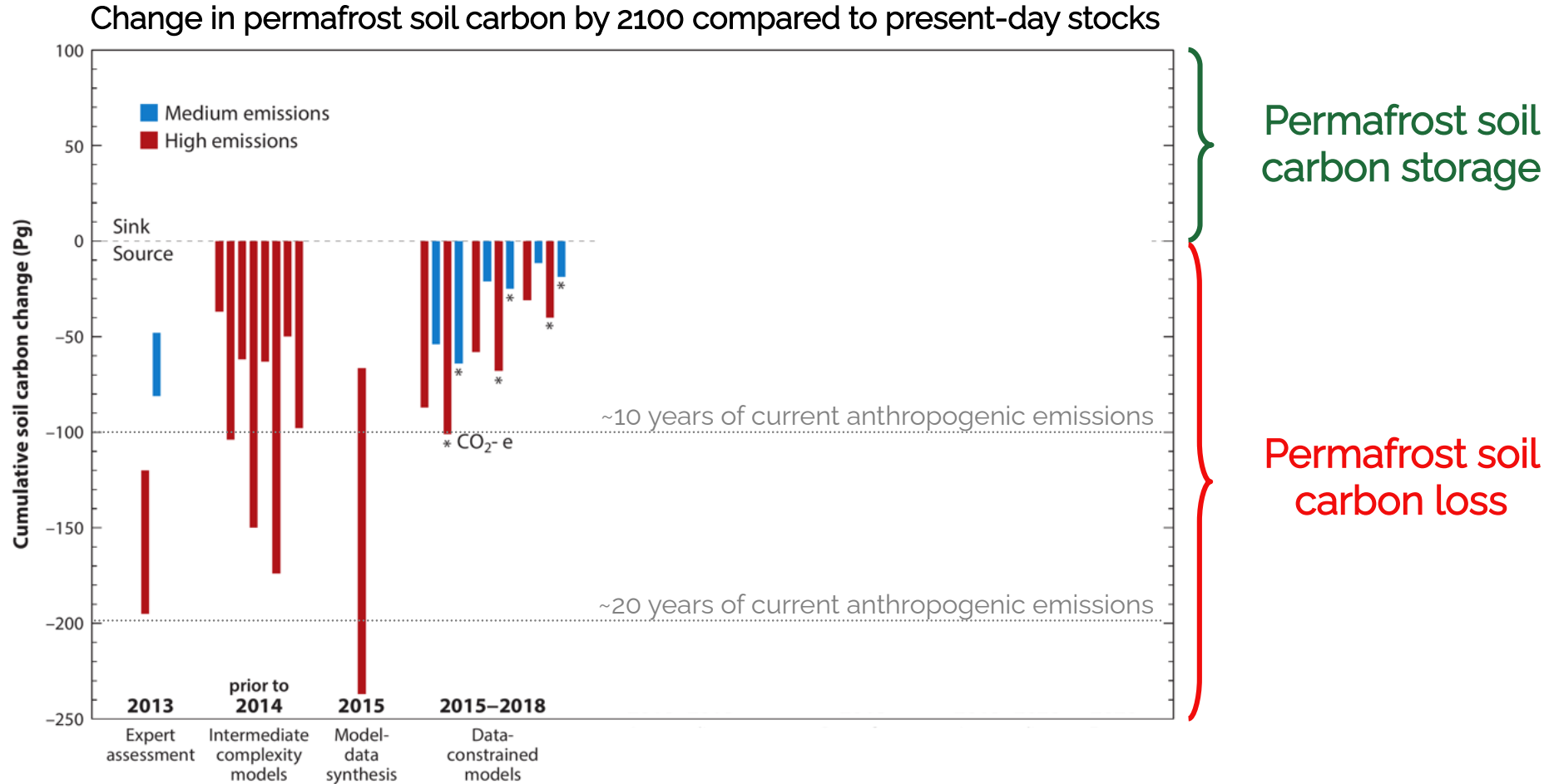
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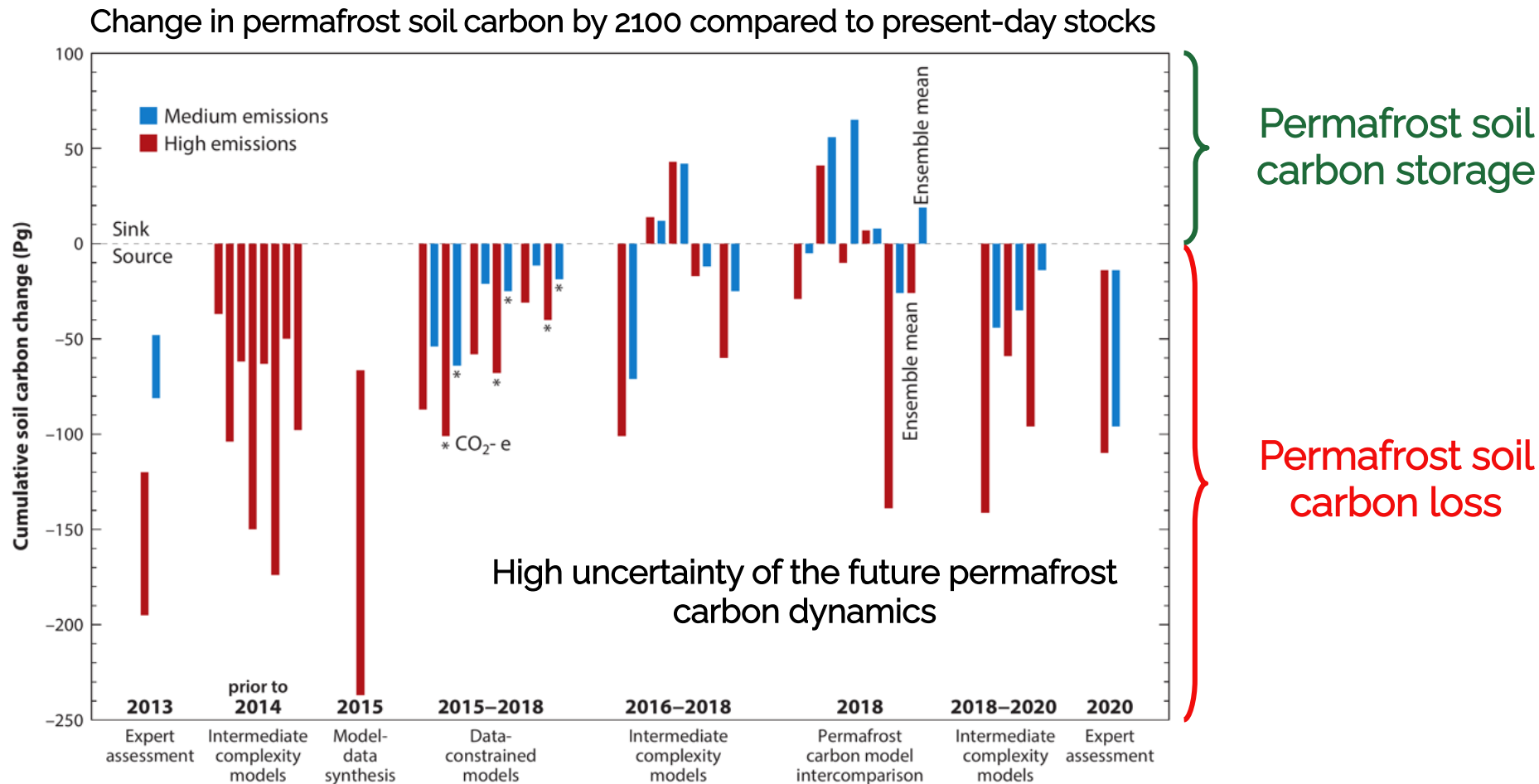
Earth system models (ESMs) or land surface models (LSMs)



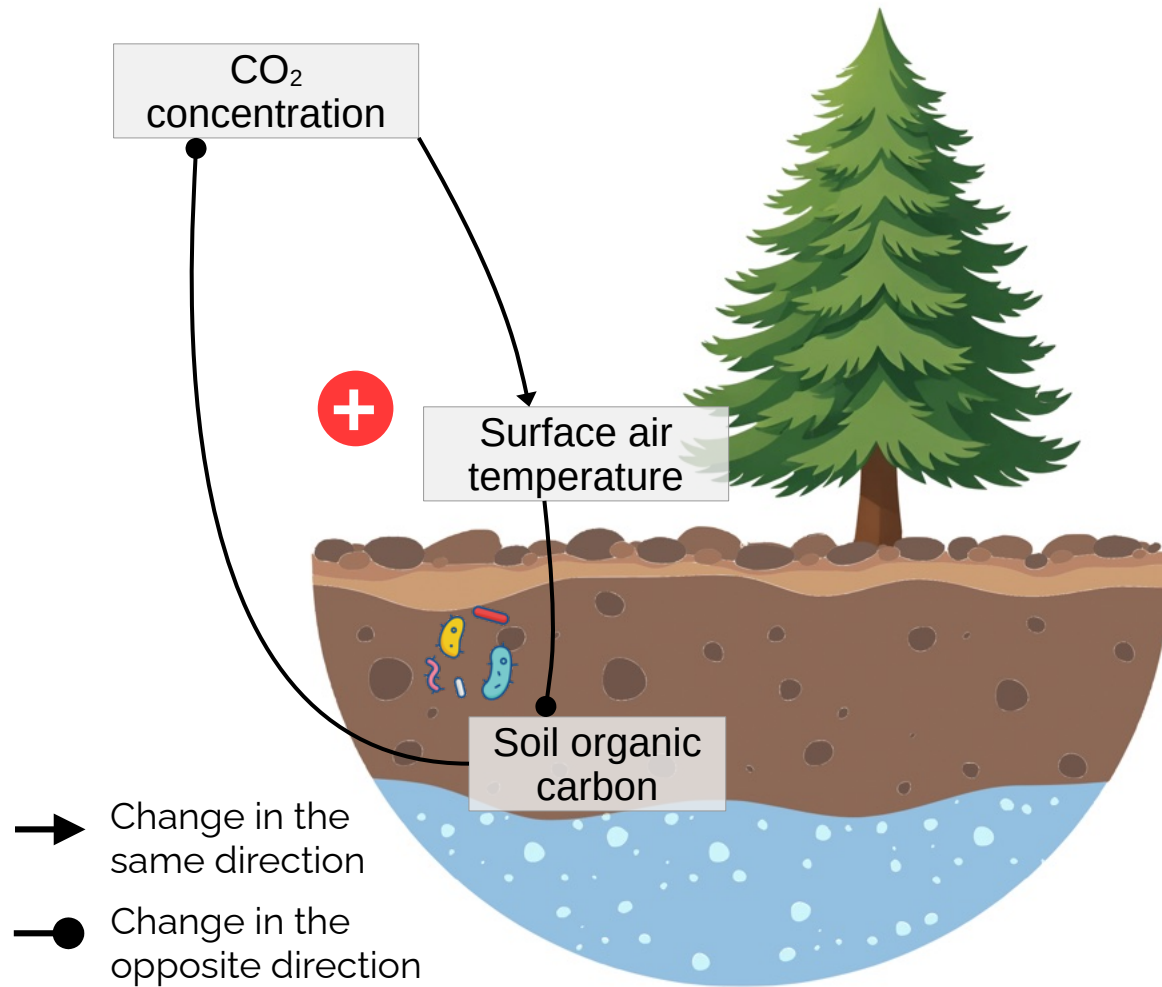
Historically, permafrost has been seen as a potential **carbon source**...



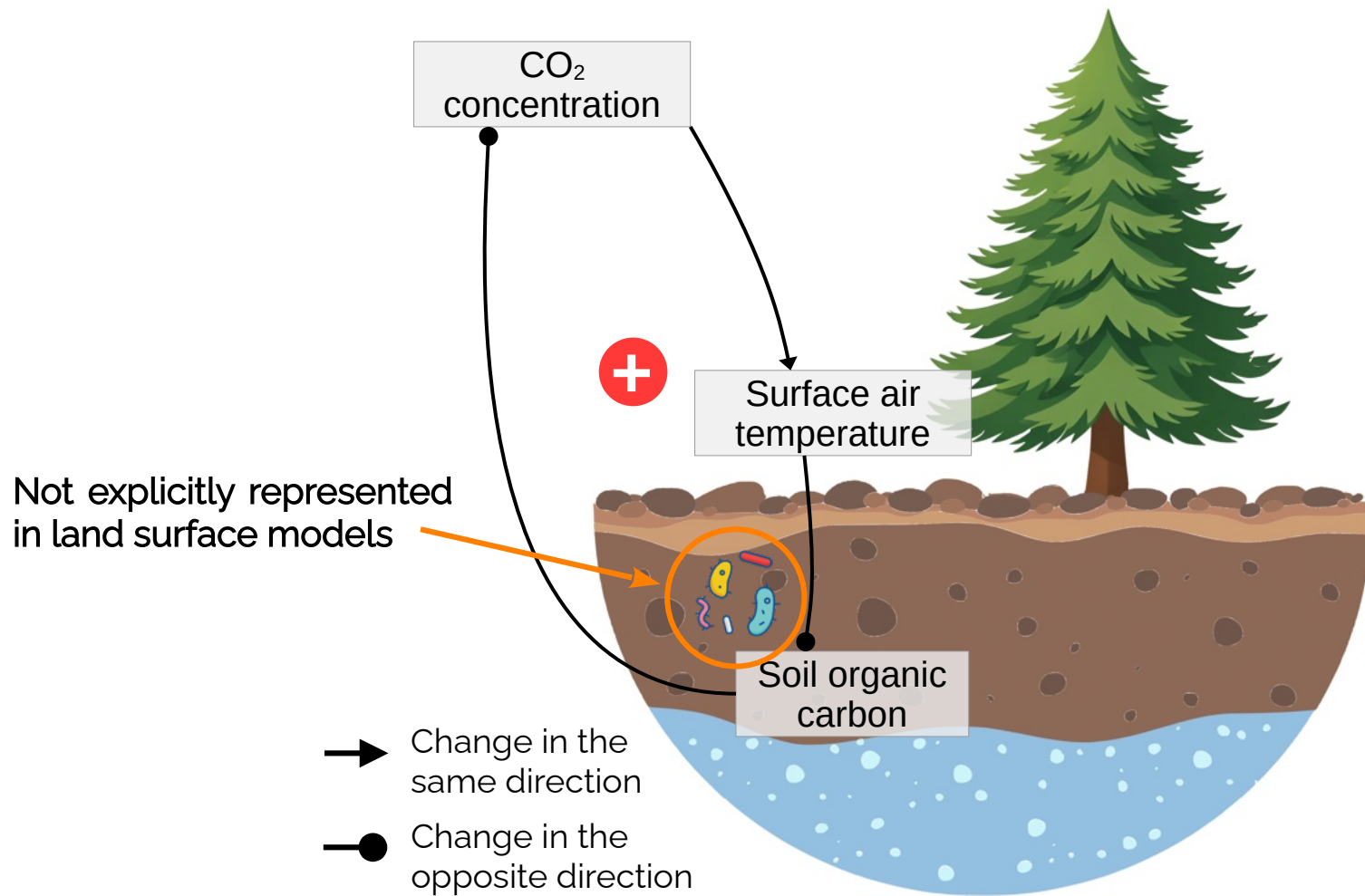
...but the story gets more complicated as models become more complex



The permafrost carbon-climate feedback depends on the **parameterization of soil respiration**



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Parameterization of soil respiration in LSMs and ESMs

→ Soil organic carbon dynamics follows a simple equation :

$$\frac{\partial SOC}{\partial t} = I - \underbrace{k \times SOC \times \theta \times \tau}_{\text{Soil respiration}}$$

SOC = soil organic carbon concentration (kgC.m^{-3})

I = SOC inputs ($\text{kgC.m}^{-3}.\text{s}^{-1}$)

k = decomposition rate (s^{-1})

θ = moisture rate modifier [0-1]

τ = **temperature rate modifier [0-1]**

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→ τ is generally represented in ESMs with :

Arrhenius equation

or

Van't Hoff laws

$$\tau = A \times e^{-Ea/RT}$$

$$\tau = Q_{10}^{(T - T_{ref})/10}$$

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Models are **highly sensitive to Q_{10}**

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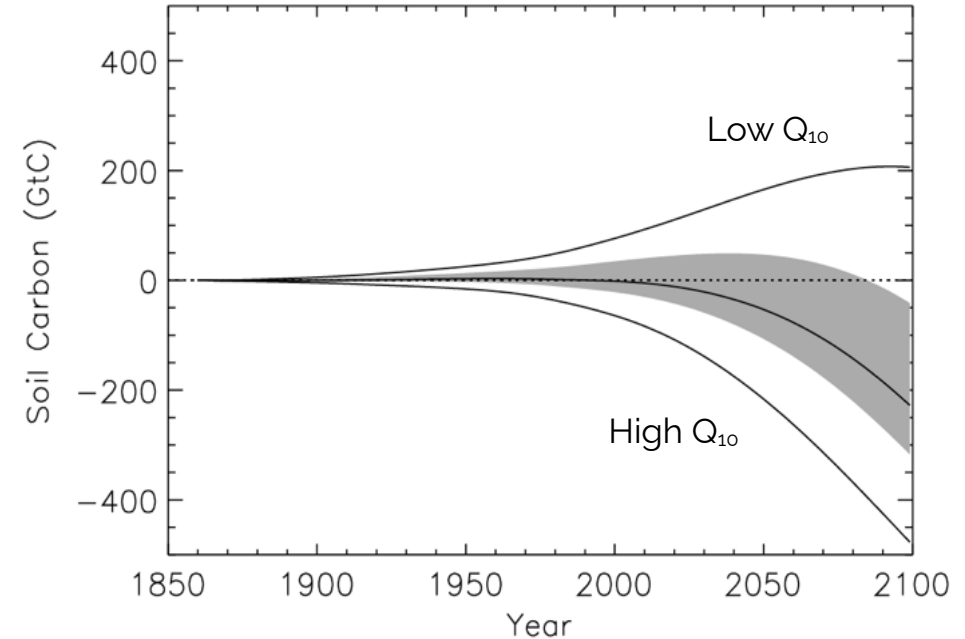
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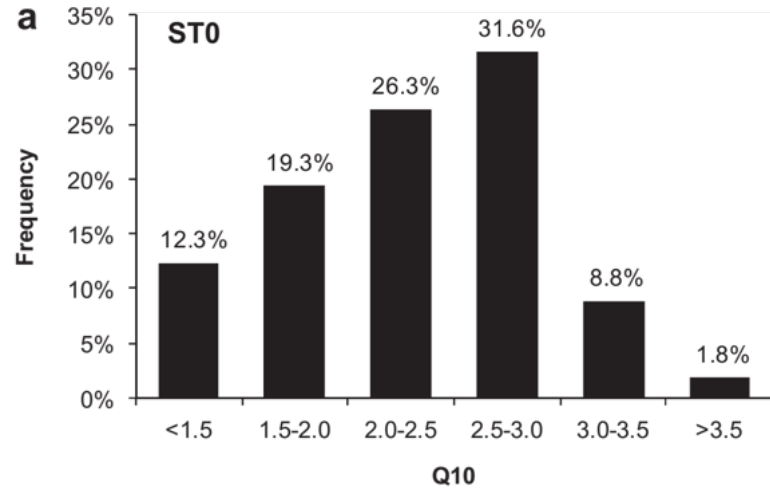
Q_{10} is generally a **fixed parameter**



→ Simple climate-carbon cycle model under a high-emission scenario

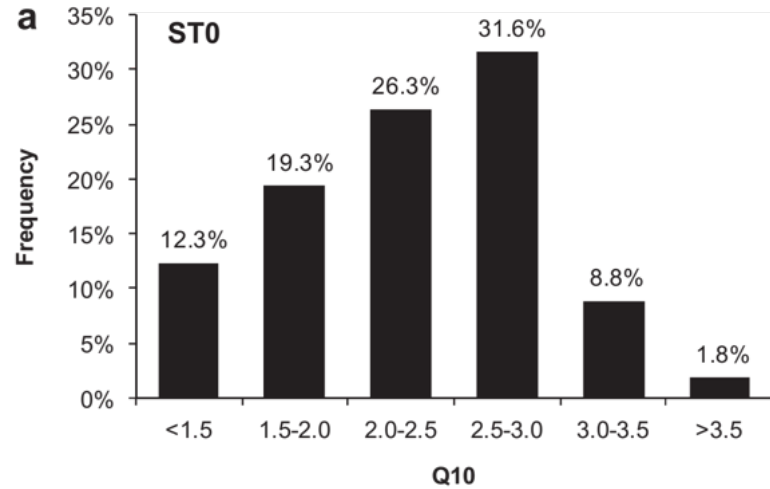
→ From +200 PgC gain to -475 PgC loss in soil carbon in 2100

But observations show that Q_{10} depends on time

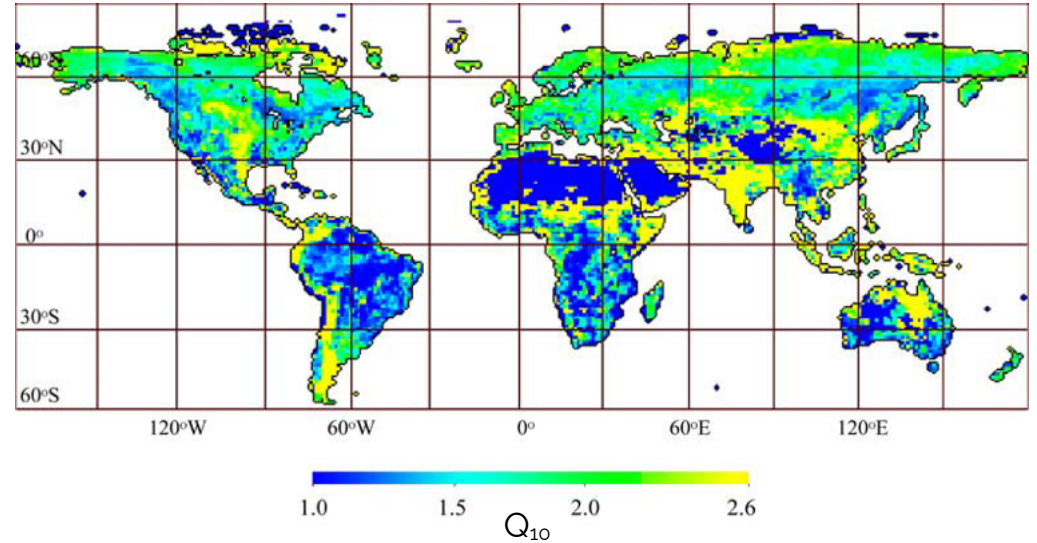


→ Seasonal variations of Q_{10} from soil field measurements

But observations show that Q_{10} depends on time and space

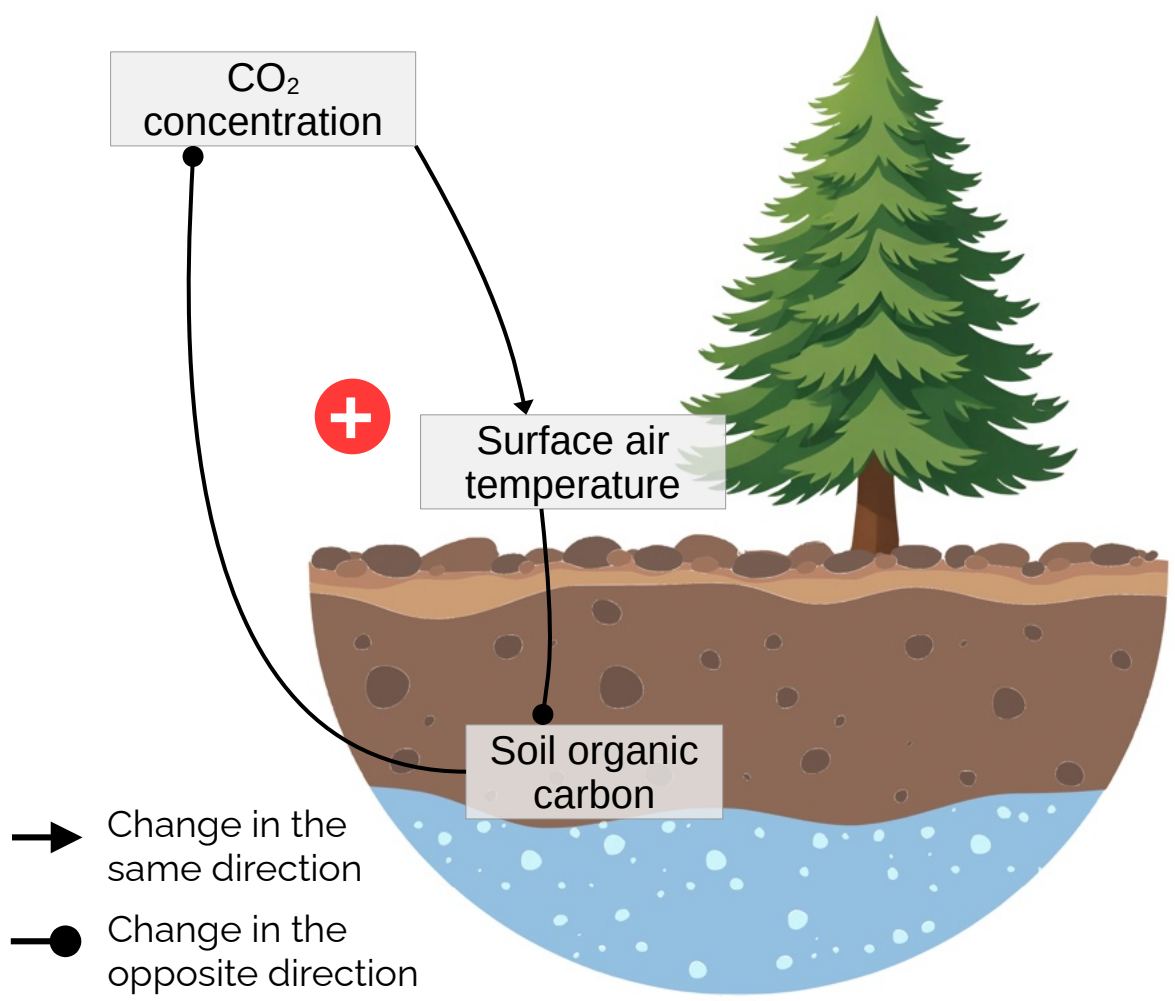


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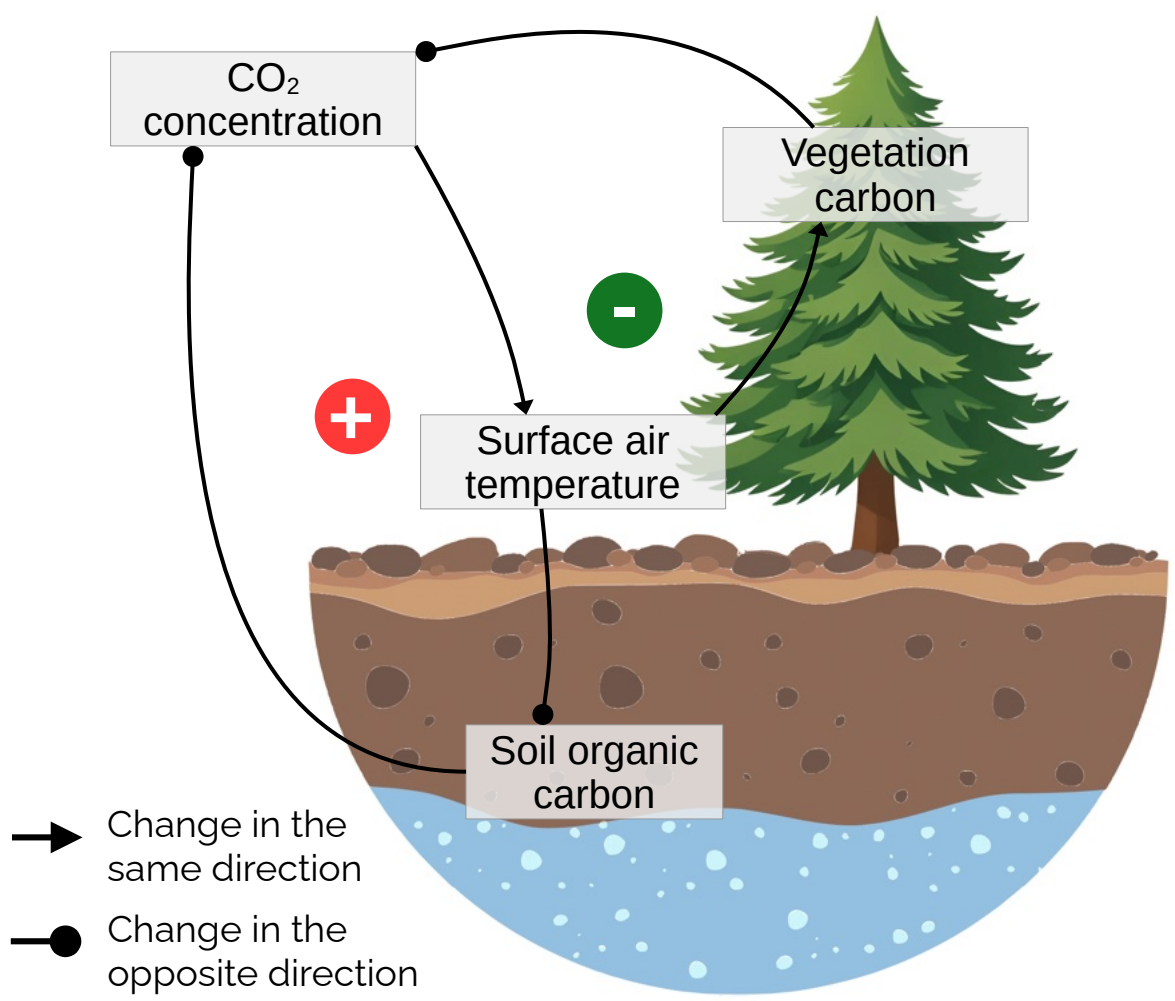


→ Spatial variations of Q_{10} from inverse modeling

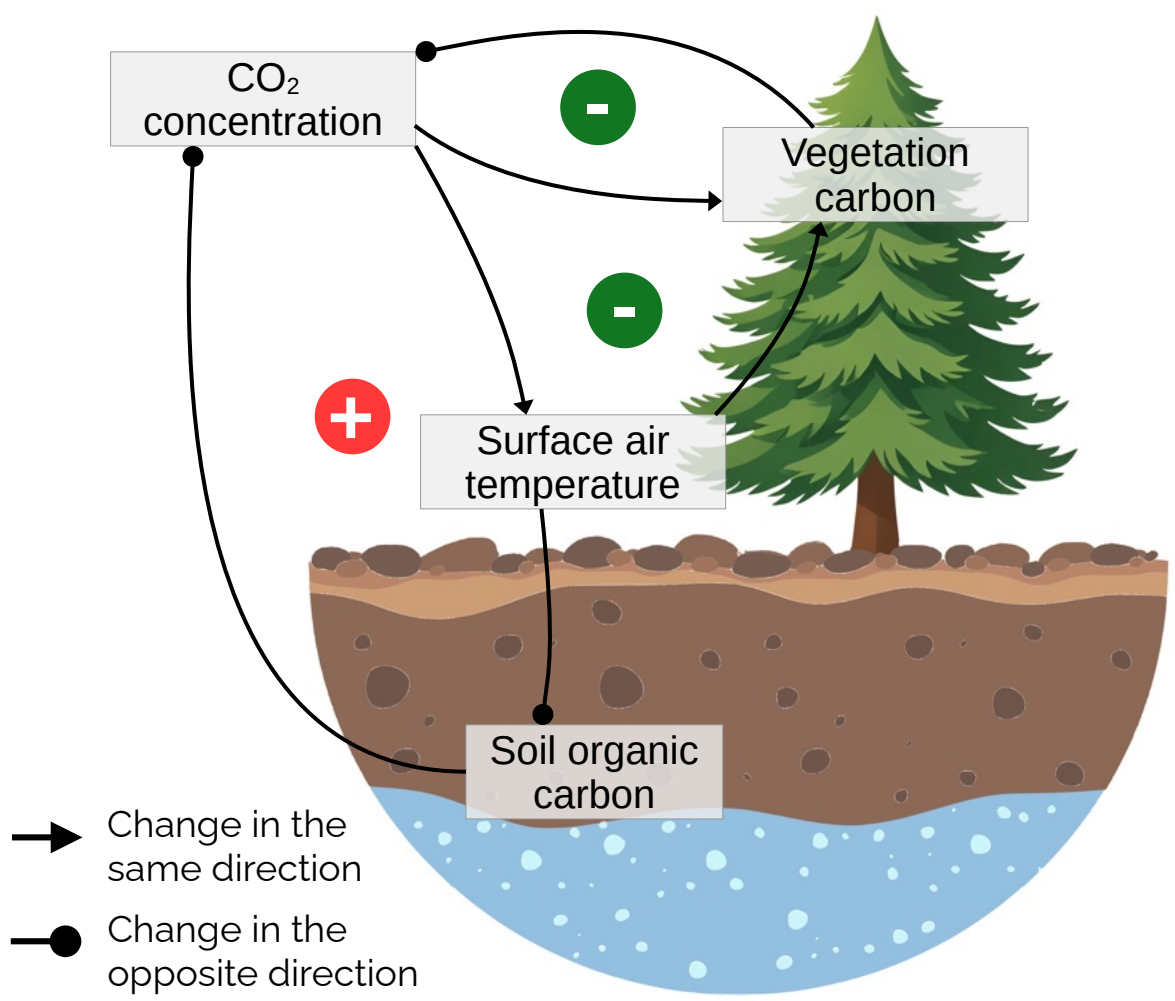
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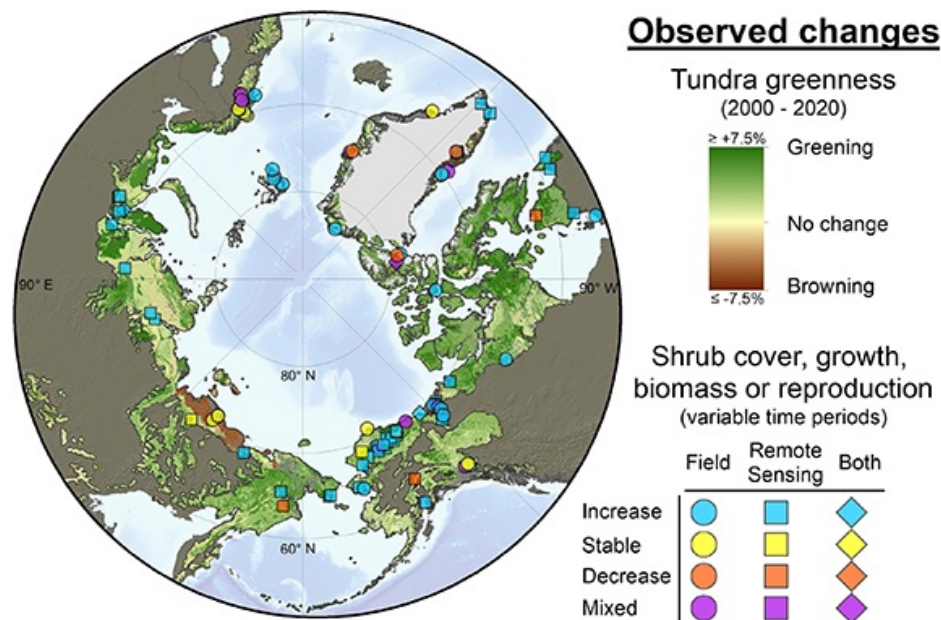


Another source of uncertainty comes from the **response of vegetation to warming and CO₂**



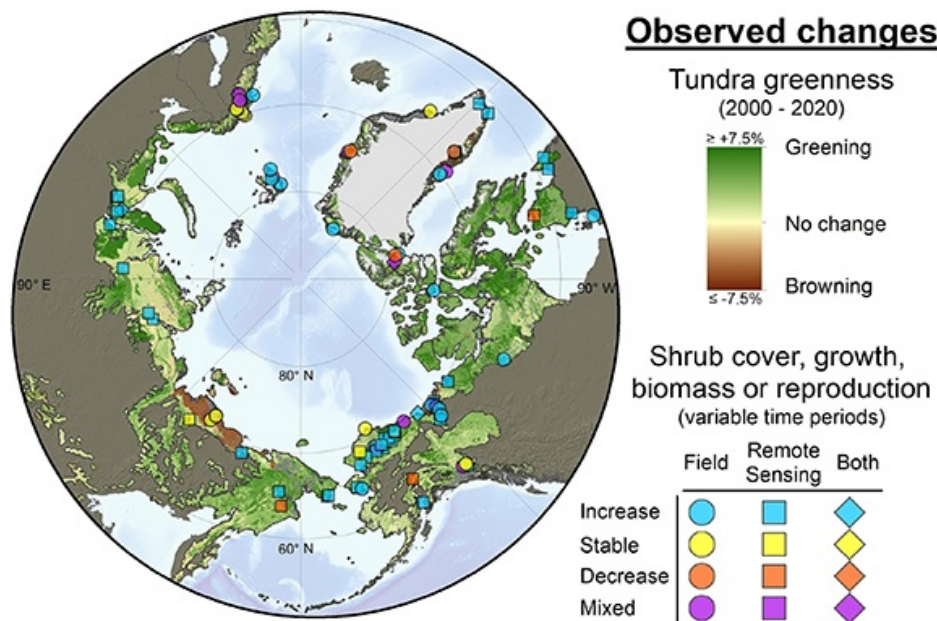
Arctic borealisation and shrubification

→ Increase in shrub biomass, growth and reproduction in the Arctic tundra.



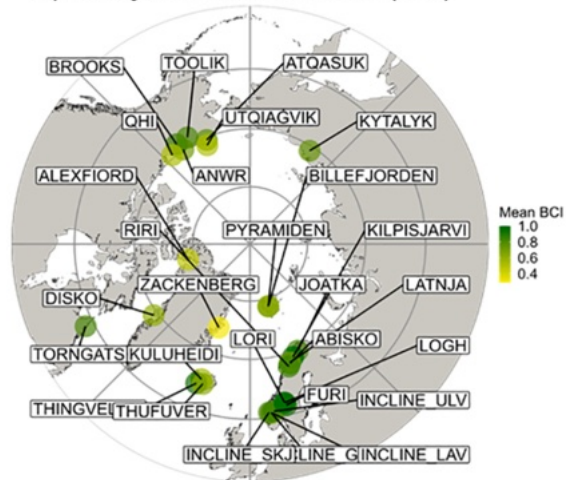
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- Increase in shrub biomass, growth and reproduction in the Arctic tundra.
- Colonisation of tundra by boreal species.

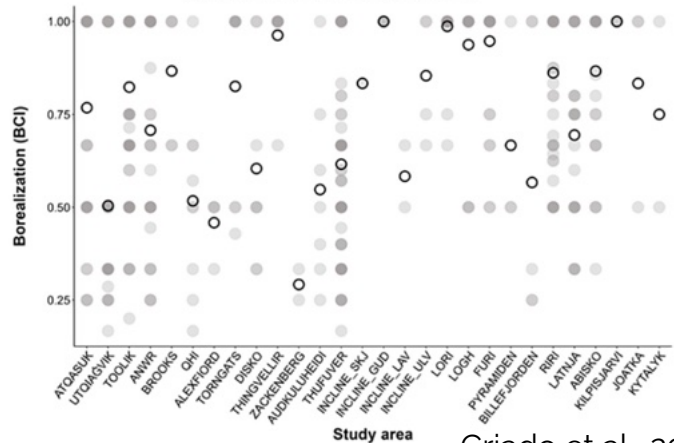


Mekonnen et al., 2021

a) Study area colonisations (BCI)



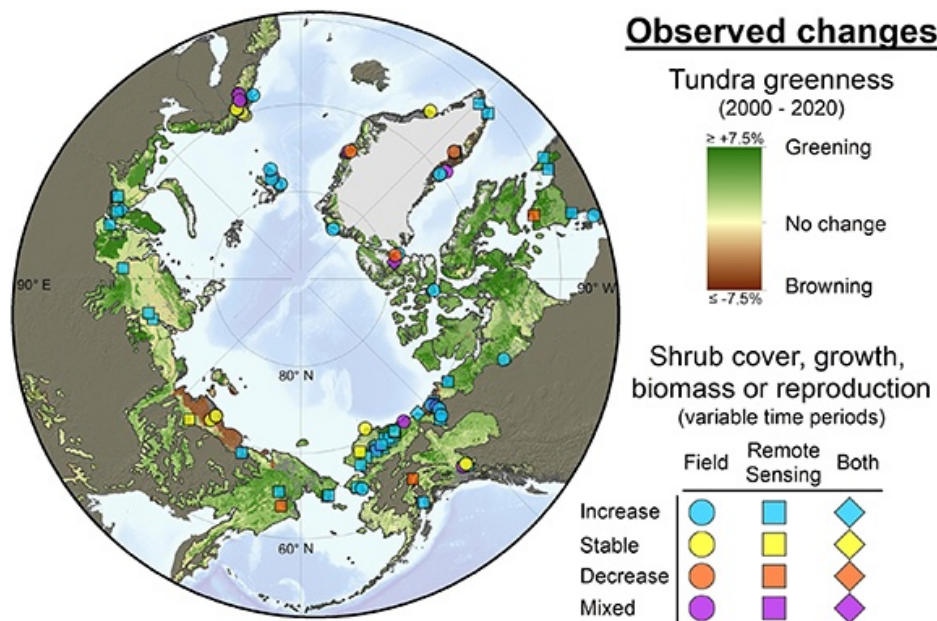
b) Plot colonisations (BCI)



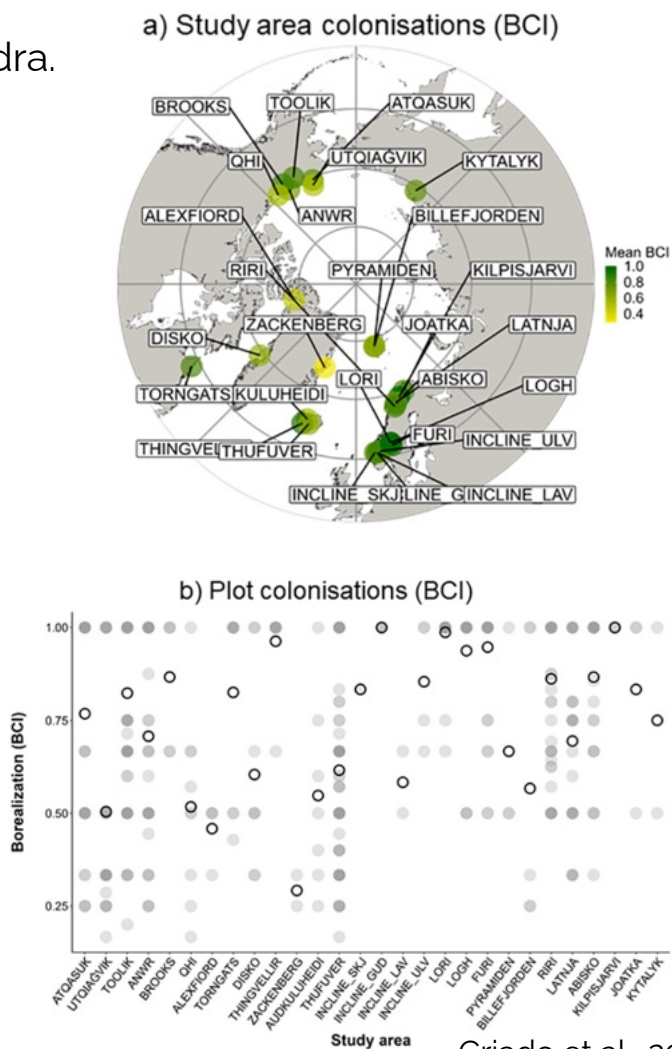
Criado et al., 2025

Arctic borealisation and shrubification

- Increase in shrub biomass, growth and reproduction in the Arctic tundra.
- Colonisation of tundra by boreal species.
- What impact on the permafrost carbon cycle ?



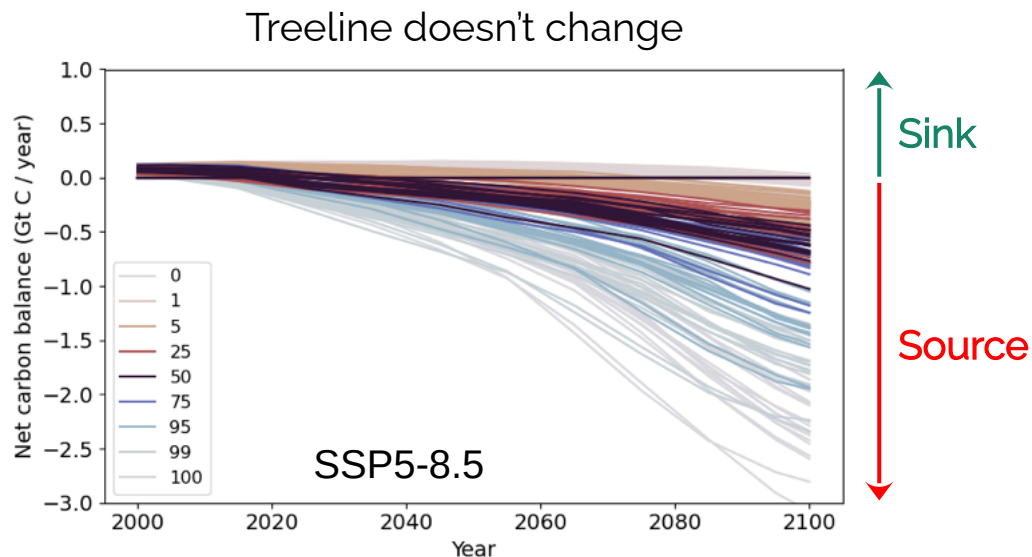
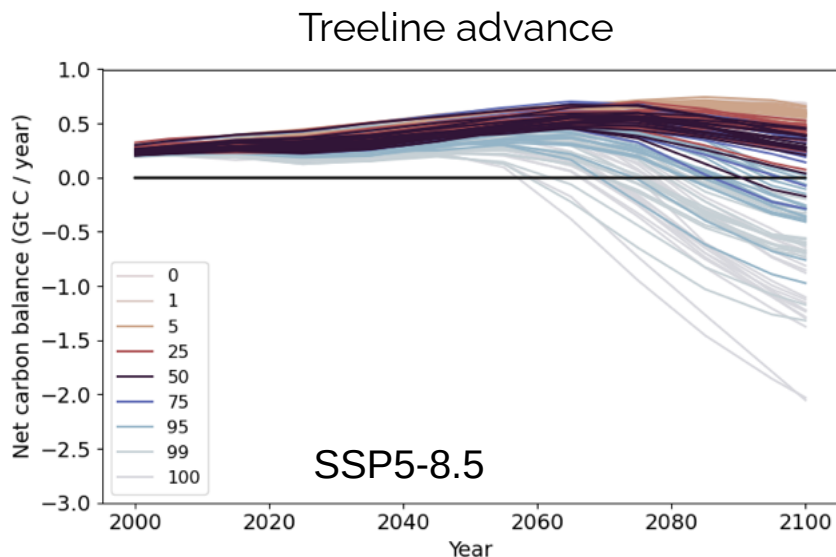
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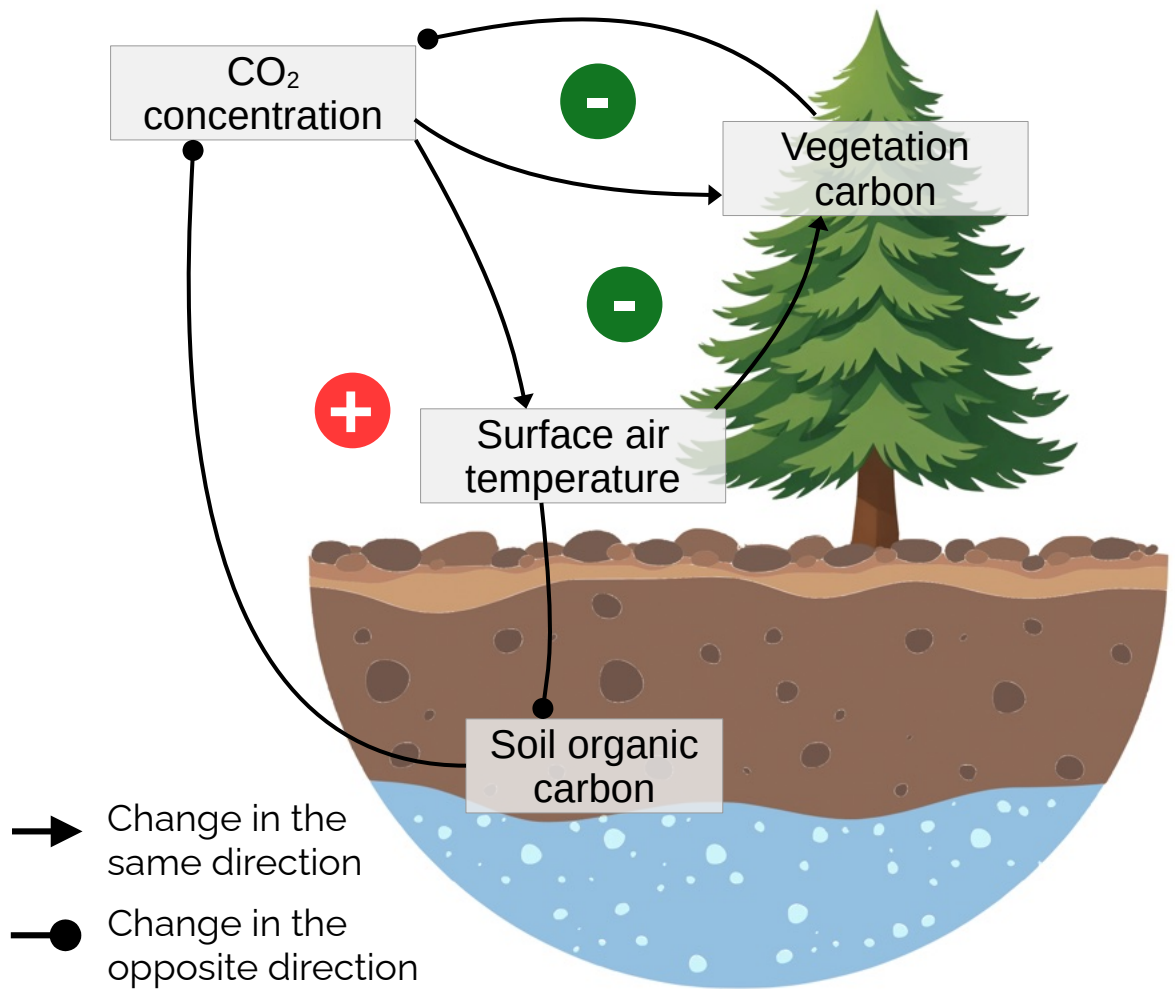
Criado et al., 2025

The **dynamics of vegetation** could have a great impact on future **carbon uptake**

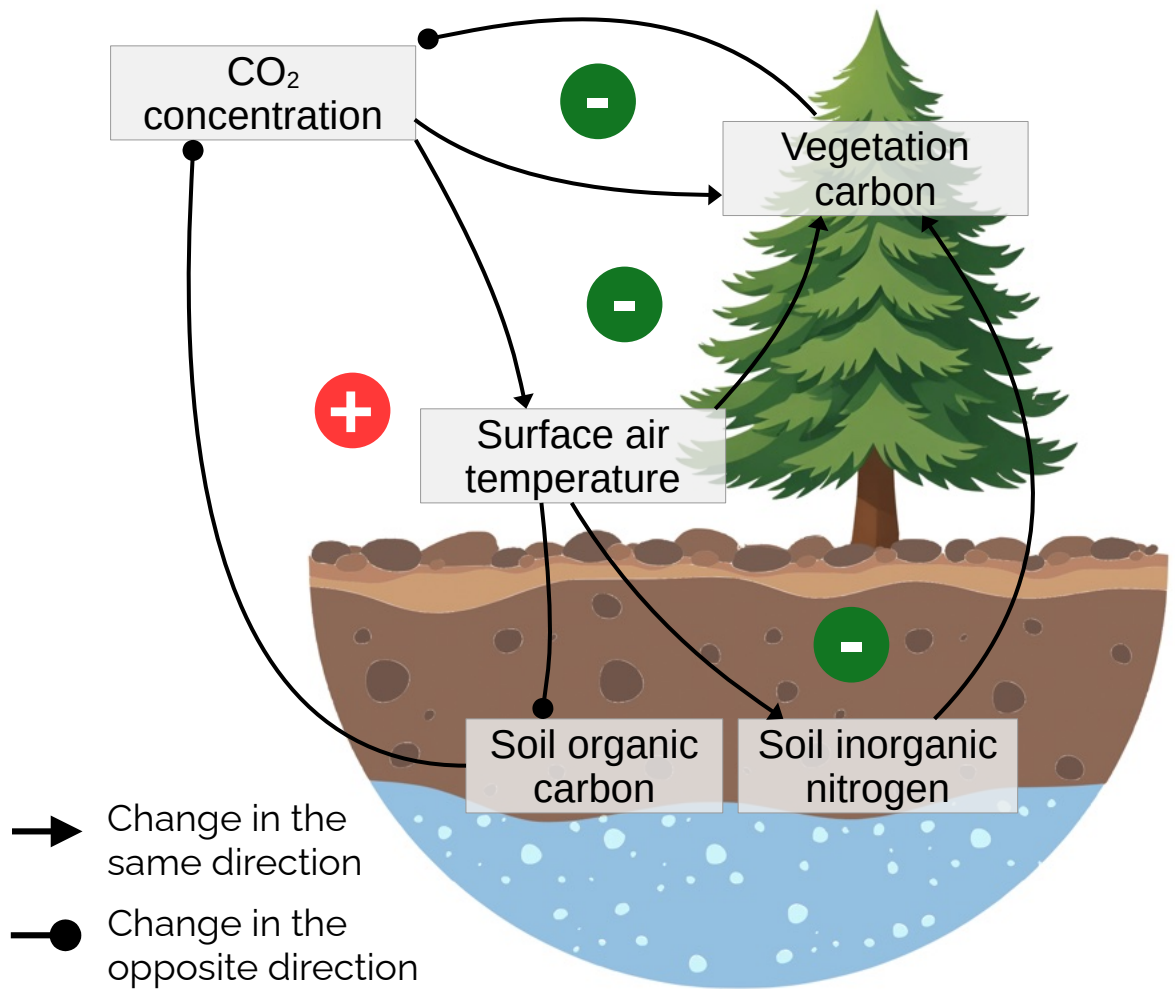
- Land surface model (JULES) simulations show a **sustained carbon uptake** due to the advance of treeline in a high-warming scenario.
- Partially offsets soil carbon losses from permafrost thaw and microbial decomposition.



Nitrogen release from permafrost thaw could also increase vegetation carbon uptake

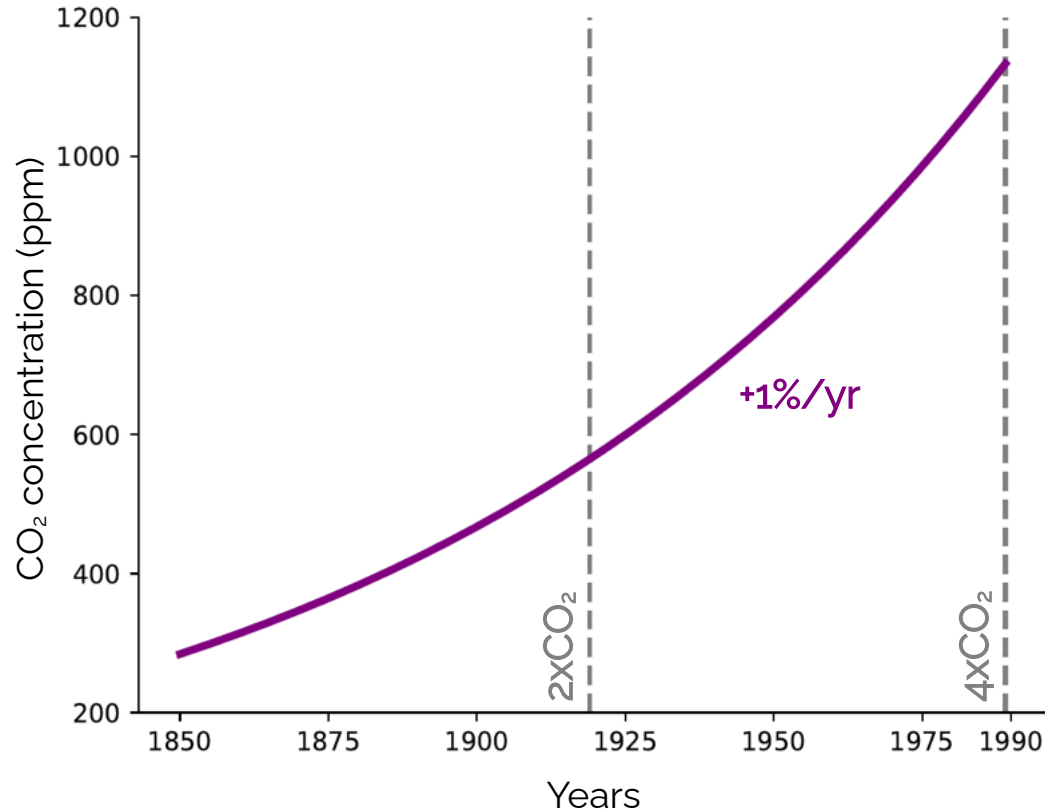


Nitrogen release from permafrost thaw could also increase vegetation carbon uptake



Idealised 1pctCO₂ experiment from C₄MIP

→ CO₂ increases at 1%/yr



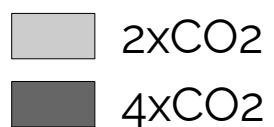
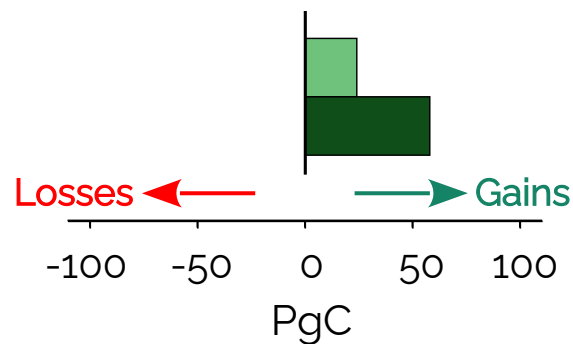
→ All other greenhouse gases at pre-industrial concentrations

→ All other forcings (land-use, nitrogen inputs...) pre-industrial

Land **gains carbon** throughout the simulation in the permafrost region

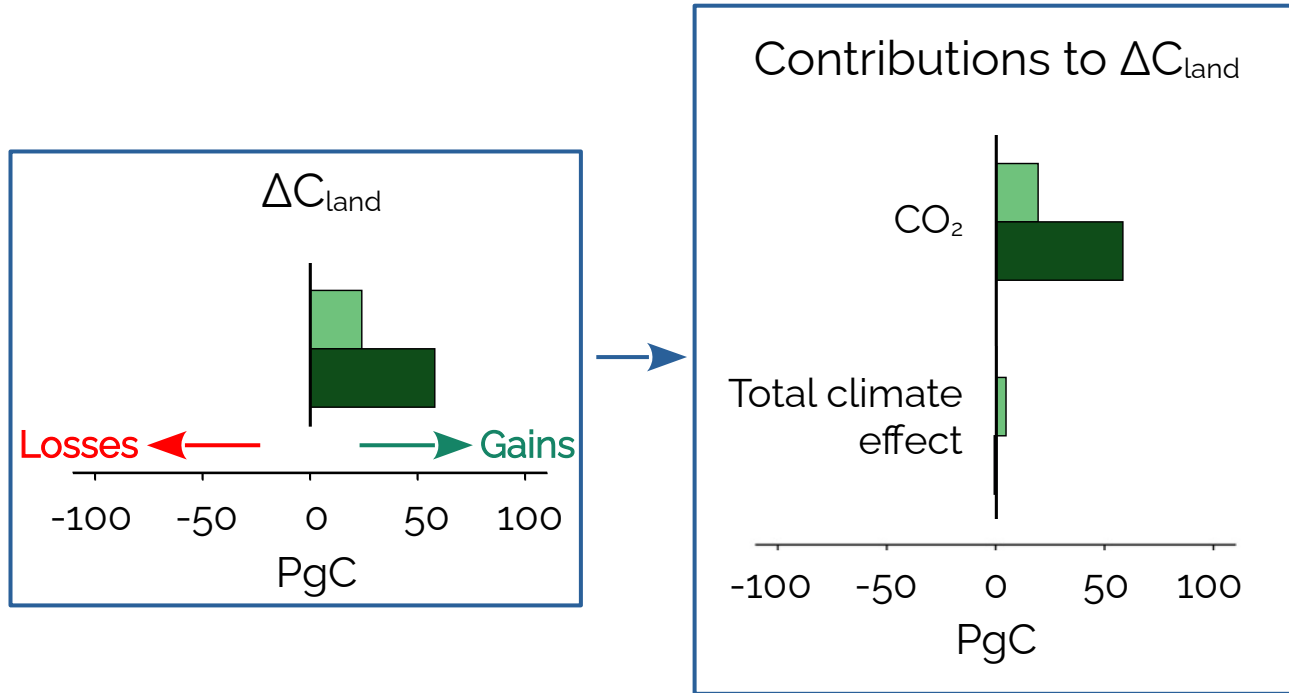
Changes in the **permafrost region** simulated by an **Earth system model** (IPSL-Perm-LandN)

ΔC_{land} = Change in total land carbon (vegetation + soils)



ΔC_{land} is driven by CO_2 fertilization

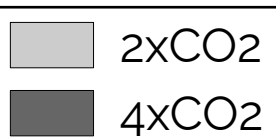
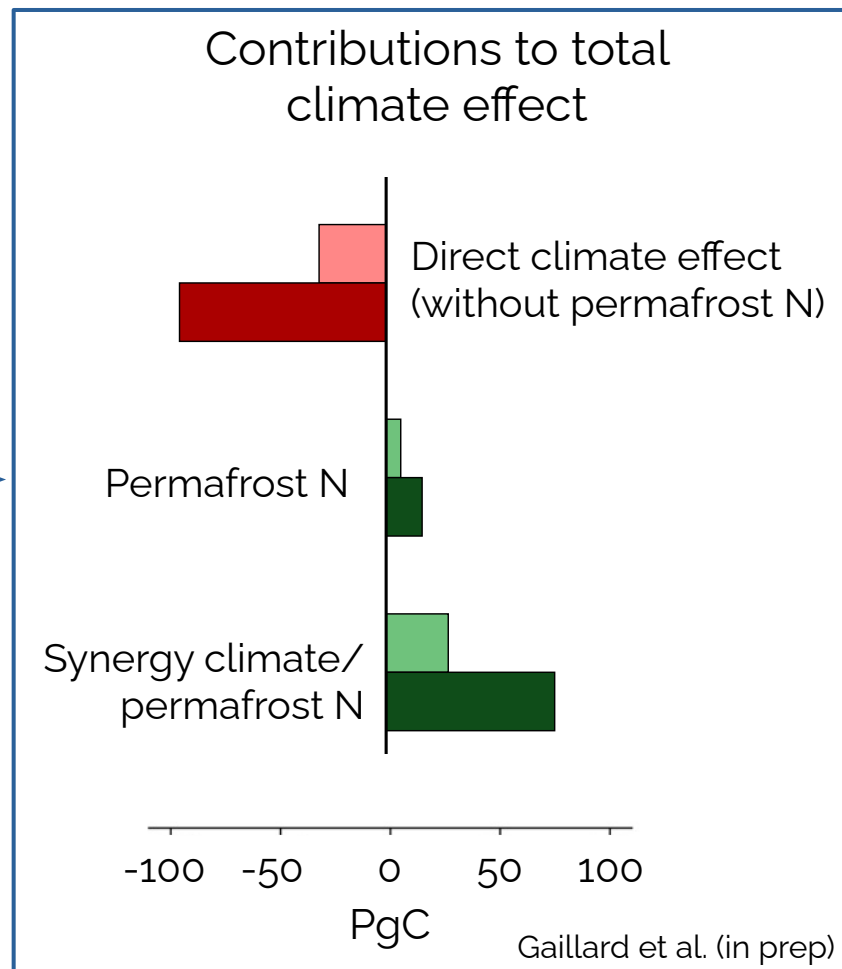
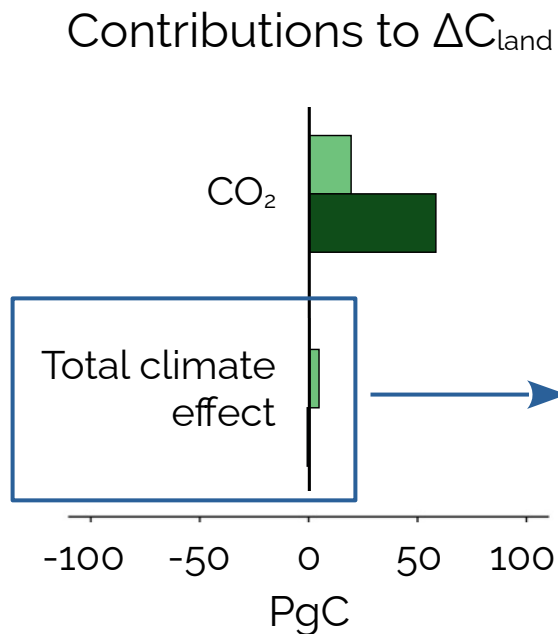
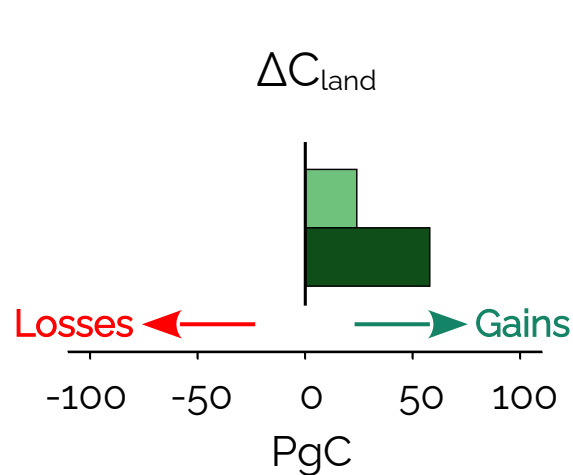
Changes in the **permafrost region** simulated by an **Earth system model** (IPSL-Perm-LandN)



→ Total climate effect is negligible

Strong nitrogen effect compensates direct climate effect

Changes in the **permafrost region** simulated by an **Earth system model** (IPSL-Perm-LandN)



Strong nitrogen effect compensates direct climate effect

Changes in the **permafrost region** simulated by an **Earth system model** (IPSL-Perm-LandN)

→ Nitrogen effect assessed in few modelling studies with **contrasting results**



Strong increase in vegetation carbon uptake due to increased nitrogen availability (Burke et al., 2022).

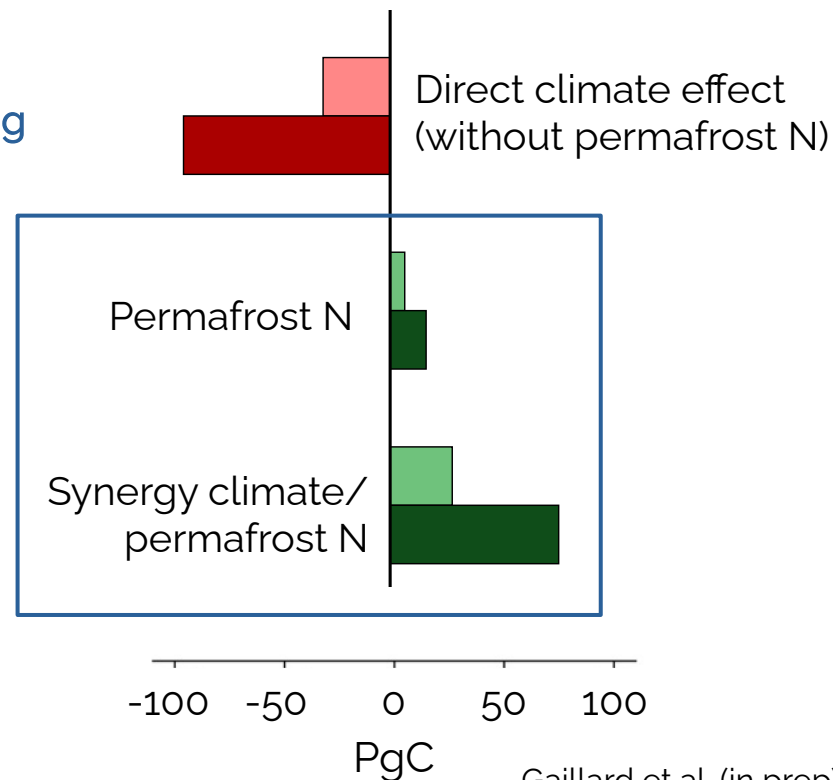


Limited to no effect of nitrogen release on vegetation carbon uptake (Lacroix et al., 2022; Koven et al., 2015).

Supporting

Contrasting

Contributions to total climate effect



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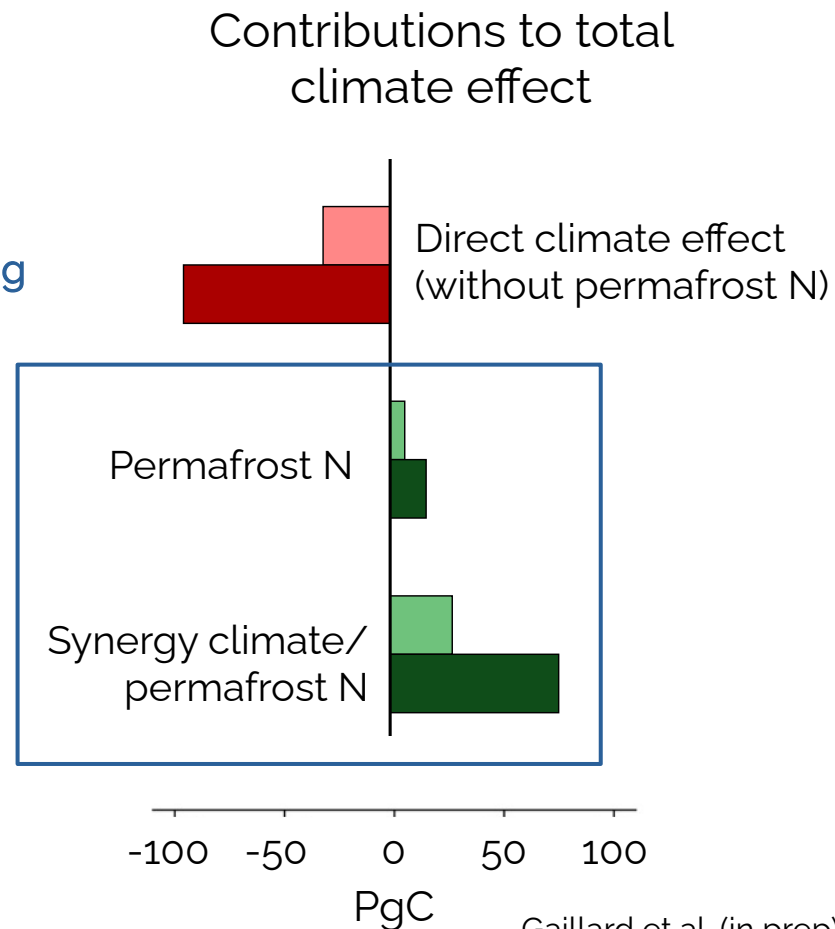
Supporting



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Contrasting

→ The representation of **nitrogen dynamics** after permafrost thaw is key for accurately predicting the **future permafrost net carbon balance**.



Conclusion

→ The future net carbon balance of the permafrost region remains **uncertain**.

→ As LSMs and ESMs become **more complex**, they start to integrate processes that could (partially?) compensate for warming-induced permafrost soil carbon losses.

→ A large part of the uncertainty originates from **biogeochemical processes**.

→ Improvements of the **physics** of models are also needed (snow dynamics, soil thermal properties and hydrology, small-scale heterogeneity, abrupt thaw processes...).

